

# Aerosols and Meteorology

## 1) Background

- 2) Aerosols and precipitation
- 3) Aerosol effects on lightning
- 4) Smoke and severe storms



# Lecture 1: Background

- Definition and examples of aerosols
- Size distributions
- Effects on radiation
- Effects on cloud physics and storm dynamics
  - Provide some observational evidence
  - Background on how aerosol-cloud interactions are modeled

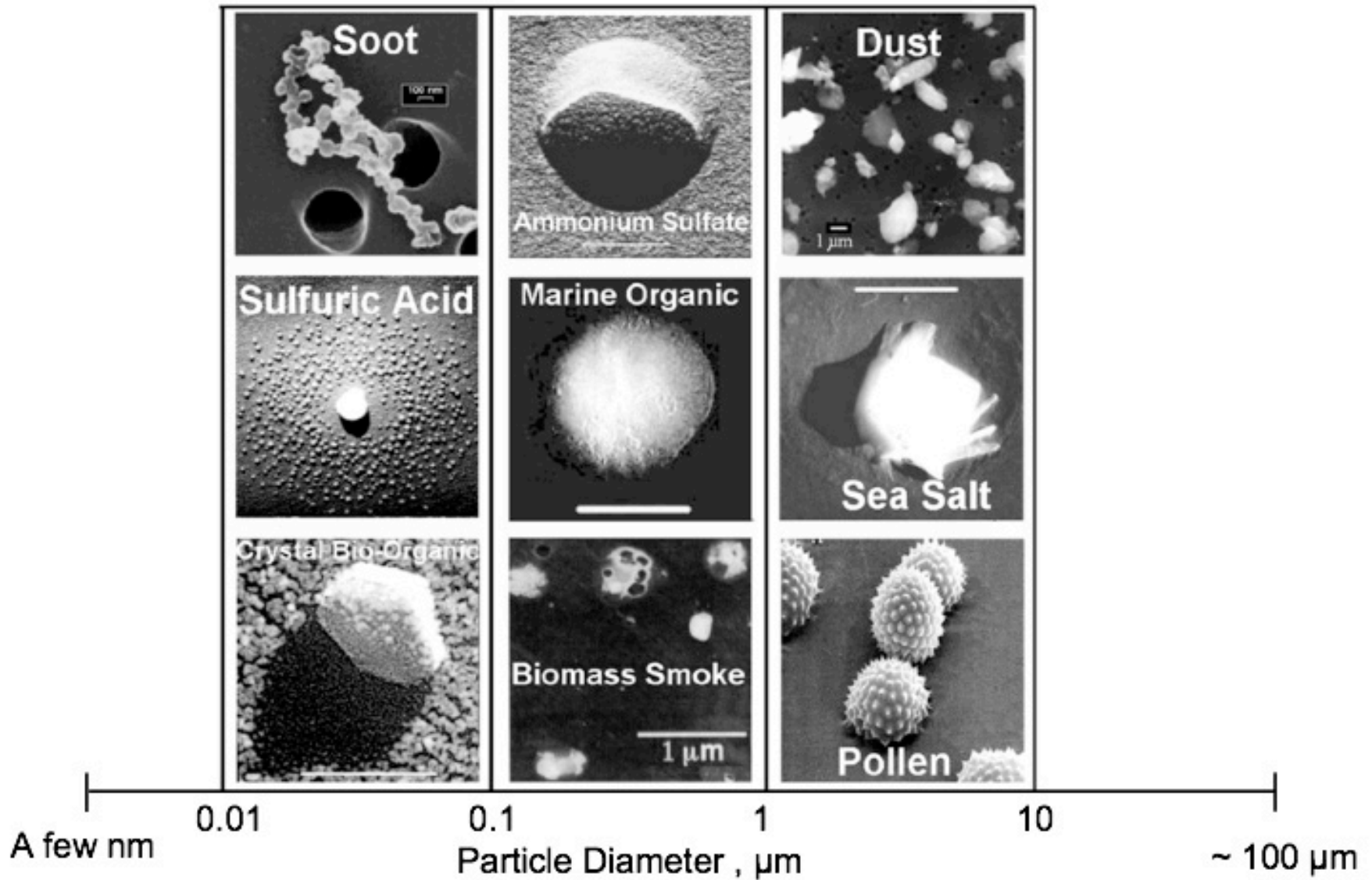
Tao et al. (2012)

Impact of Aerosols on Convective Clouds and Precipitation, *Rev. Geophys.*

# What are aerosols?

- An aerosol is defined as a colloidal system of solid or liquid particles in a gas. An aerosol includes both the particles and the suspending gas, which is usually air.

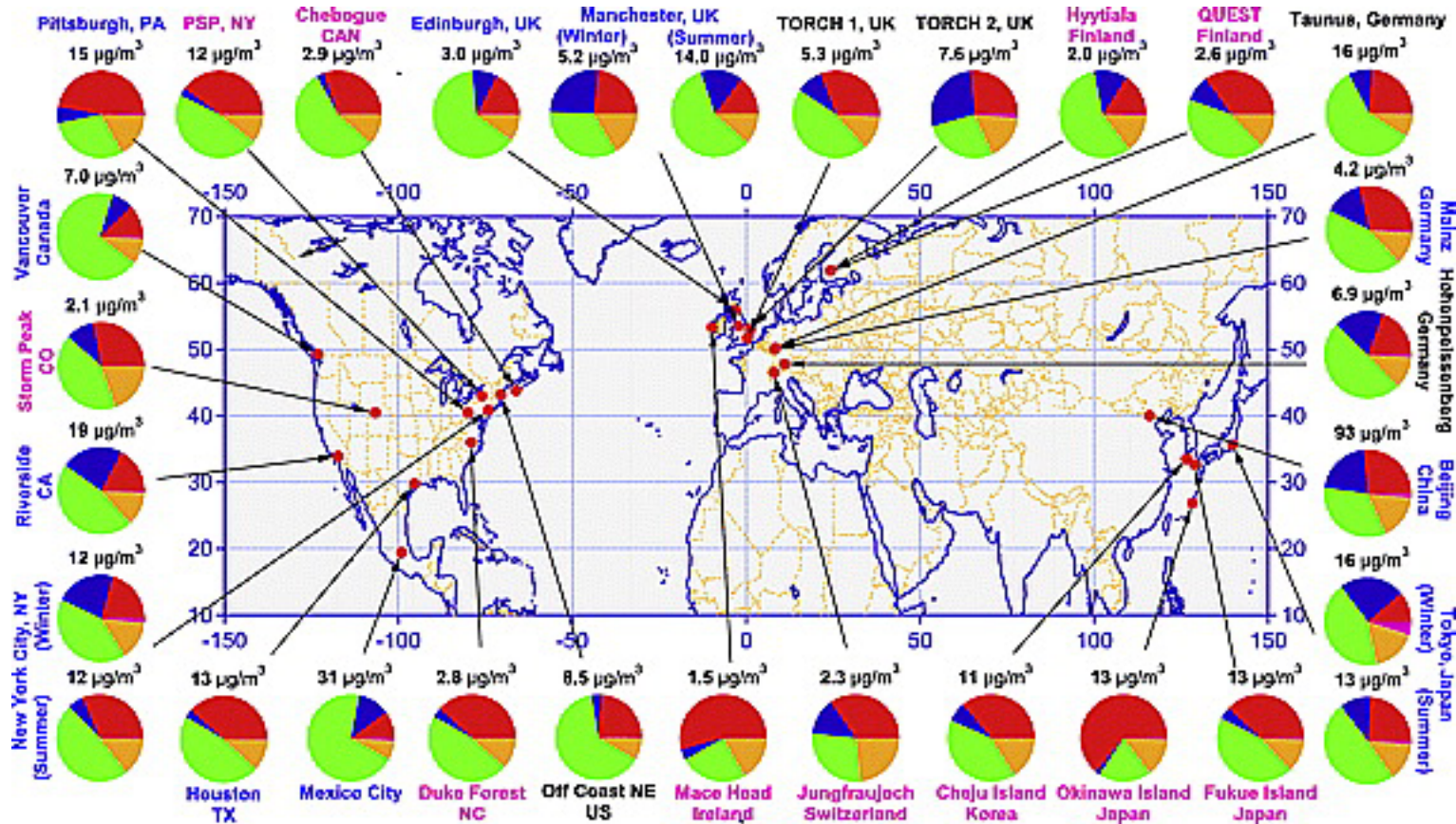
# What are aerosols?





# What are aerosols?

## Aerosol Composition at Urban and Rural Locations (Zhang et al., 2007)



Sulfate, nitrate, ammonium, organics

Not included: black carbon, dust, sea salt, water

Measurements from Aerosol Mass Spectrometer which analyzes aerosols  $< 1000 \text{ nm}$

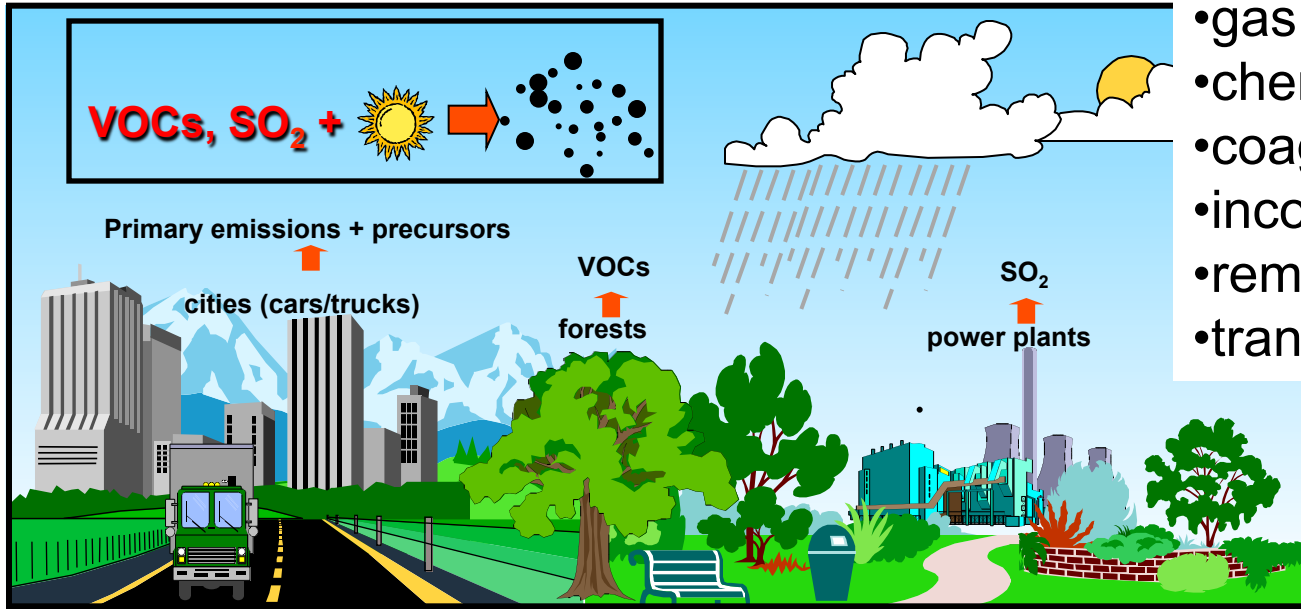
# What are aerosols?

- An aerosol is defined as a colloidal system of solid or liquid particles in a gas. An aerosol includes both the particles and the suspending gas, which is usually air.
- *Primary aerosols* contain particles introduced directly into the gas;
- *Secondary aerosols* form through gas-to-particle conversion.



## ***Secondary (produced in atmosphere) sources of particles:***

- vehicles (organics, nitrates)
- industry (sulfate, organics, nitrates, ammonium)
- plants (organics)
- agriculture (ammonium, nitrate)



## ***Atmospheric processes:***

- gas  $\rightarrow$  particle conversion
- chemical reactions
- coagulation
- incorporation into clouds
- removal by precipitation
- transport

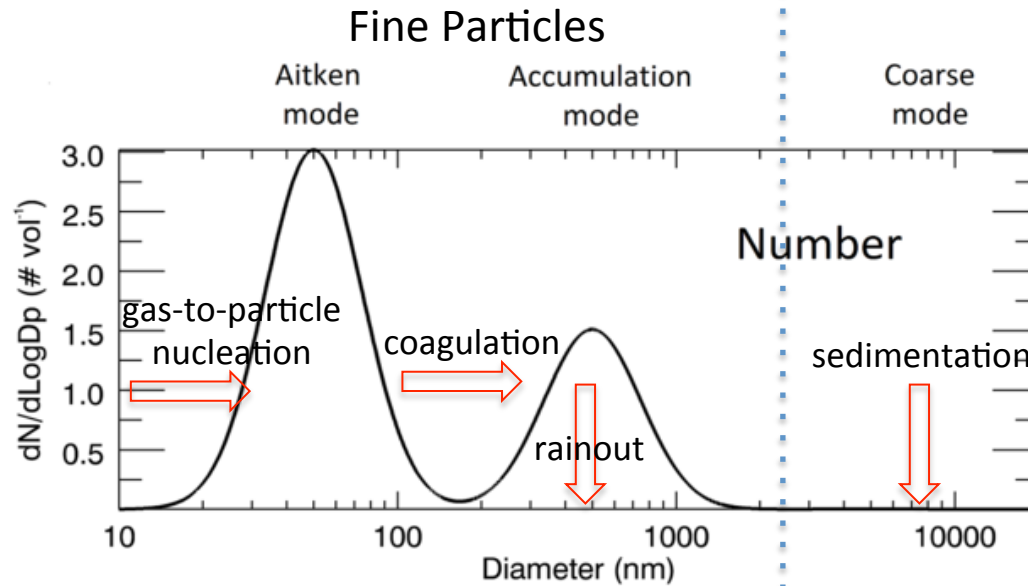
## ***Primary (directly emitted) sources of particles:***

- vehicles (soot, organics)
- industry (soot, sulfate, organics, metals)
- construction & agriculture (soot, soil)
- sea-spray (salt)
- fires (soot and organics)

# Aerosol Size Distribution

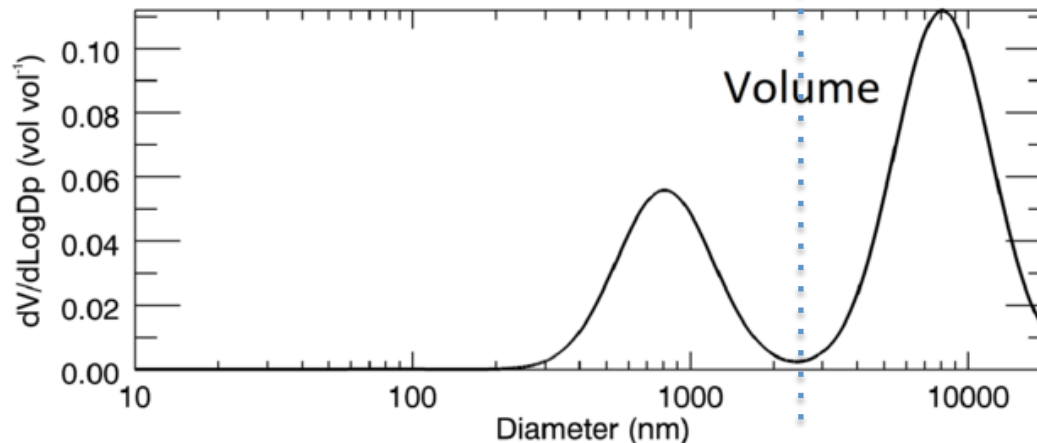
## Chemical Processes

Sulfate  
Nitrate  
Ammonium  
Organics  
Black carbon



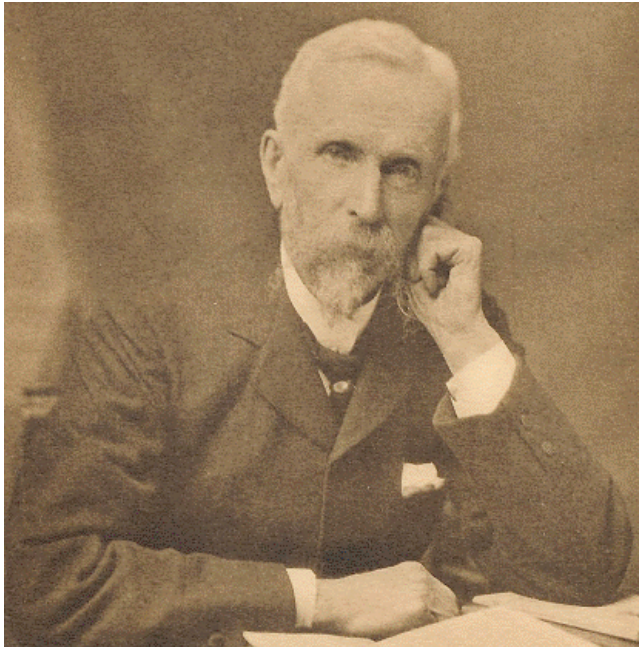
## Mechanical Processes

Dust  
Sea salt  
Pollen  
Spores  
Ash (volcanos)





**“We have in this fine dust a most beautiful illustration of how the little things in the world work great effects by virtue of their numbers.”**



-John Aitken, 1880

<http://www.iara.org/AerosolPioneers.htm>

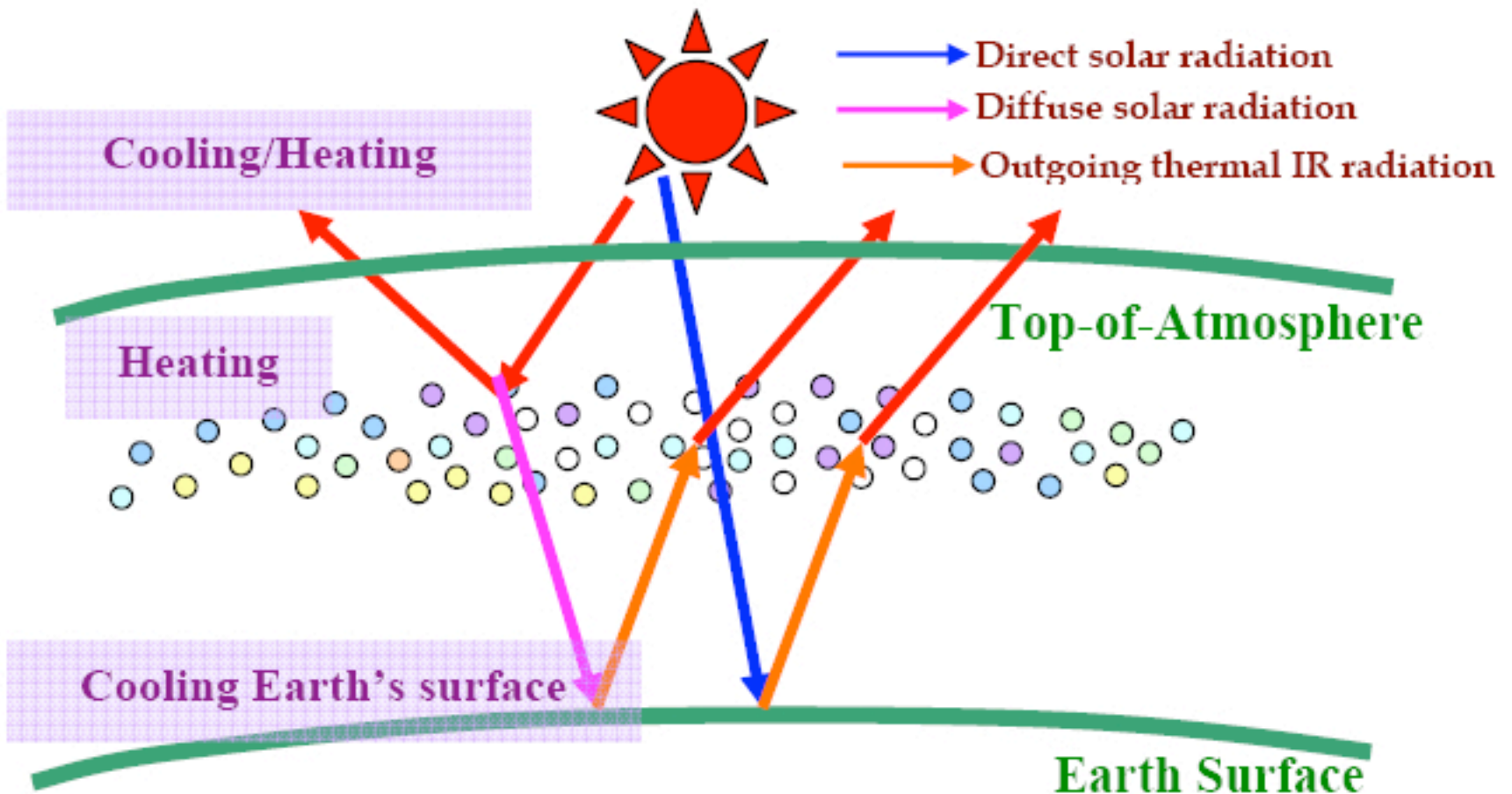
# Why are aerosols important?

- Aerosols affect the energy budget of the atmosphere by scattering and/or absorbing radiation
- Aerosols play a key role in forming cloud droplets and ice crystals
- Aerosols affect cloud properties, e.g. albedo, lifetime, precipitation



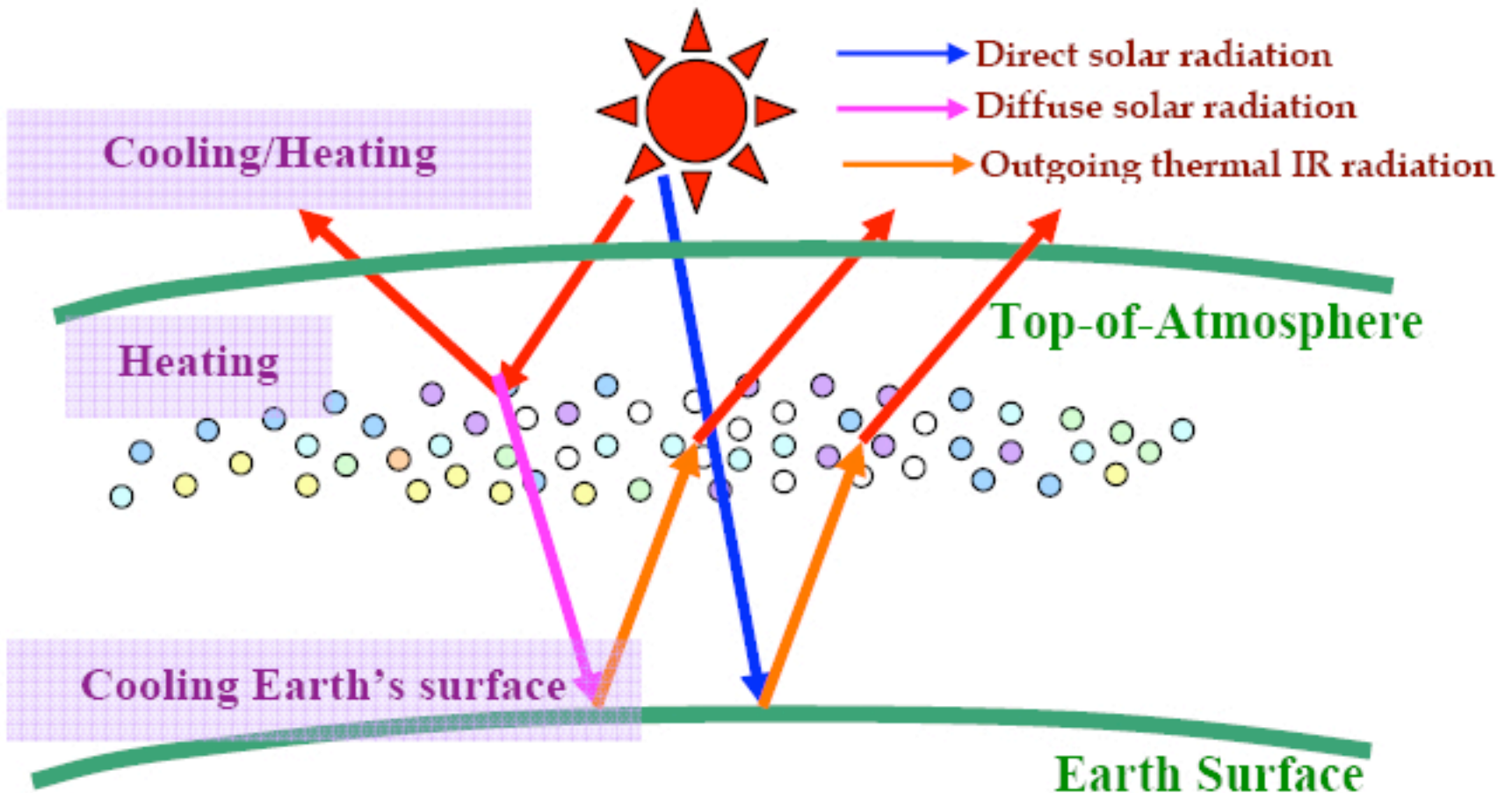
# Aerosols and Radiation

- Aerosols scatter solar radiation cooling the atmosphere
- Aerosols absorb solar radiation heating regions of the atmosphere

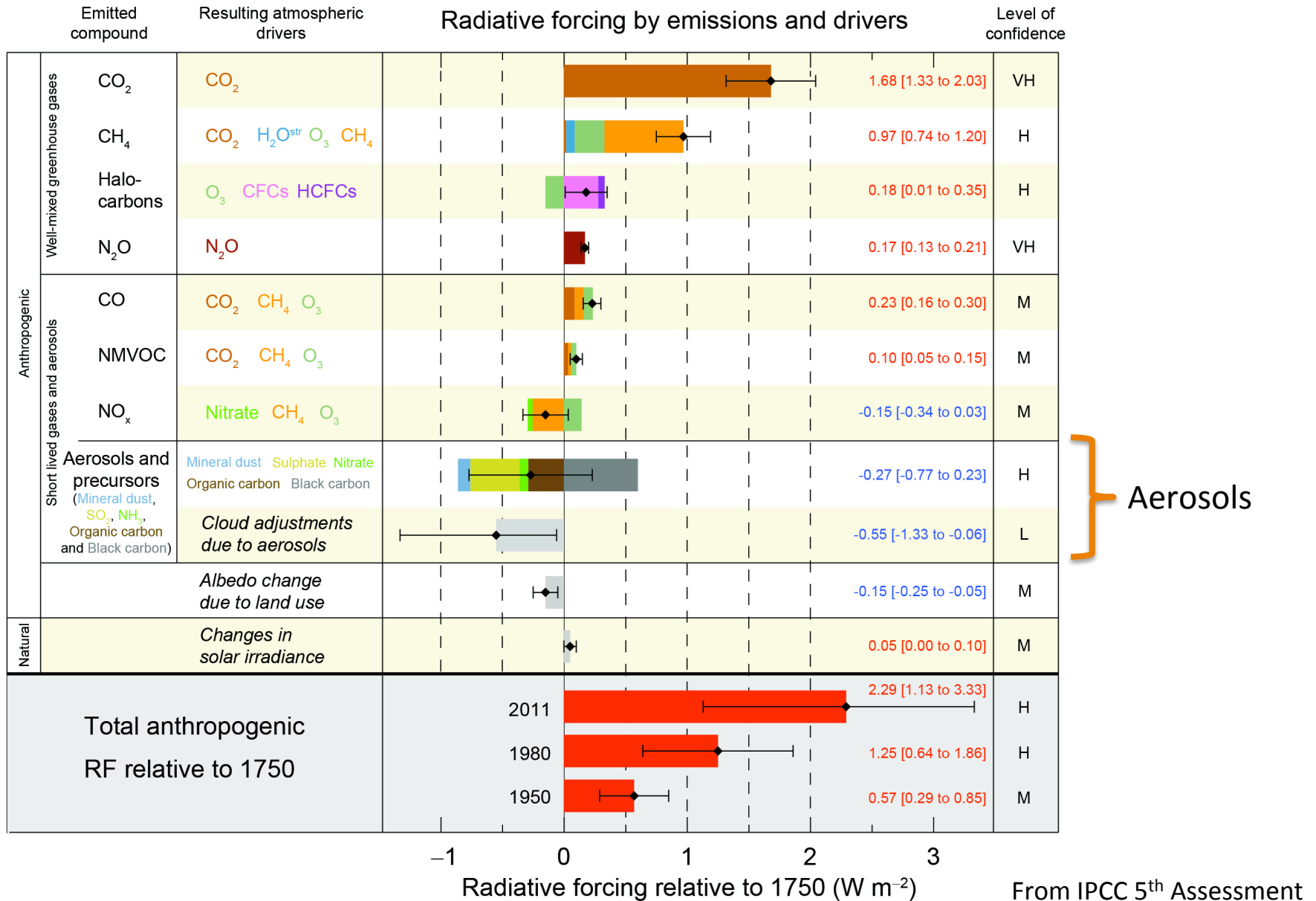


# Aerosols and Radiation

- Aerosols scatter solar radiation cooling the atmosphere (**sulfate, nitrate, sea salt**)
- Aerosols absorb solar radiation heating regions of the atmosphere (**BC, dust, BrC**)

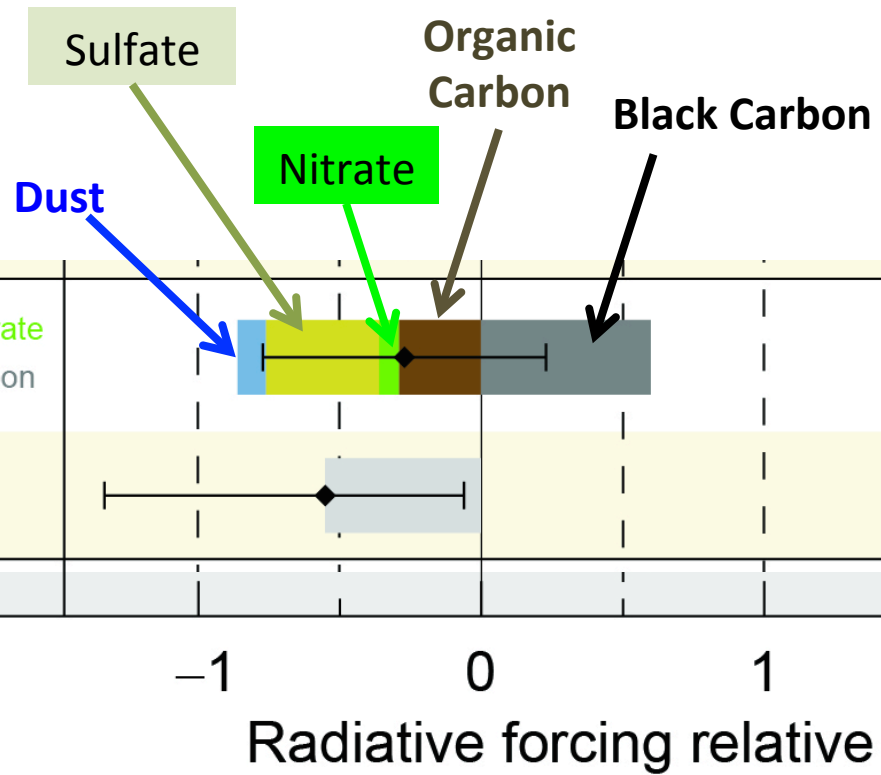


# Aerosols and Climate





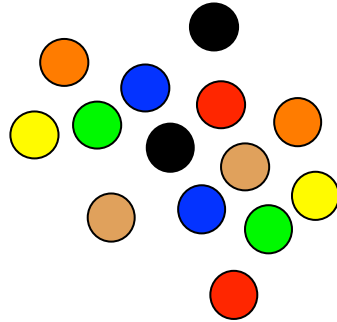
# Aerosols and Climate



# Aerosol Mixtures

## External Mixture:

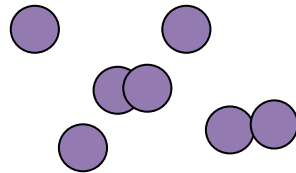
Each particle is separate in composition: BC, sulfate, sea salt



No combinations, e.g. sulfate coating BC aerosol

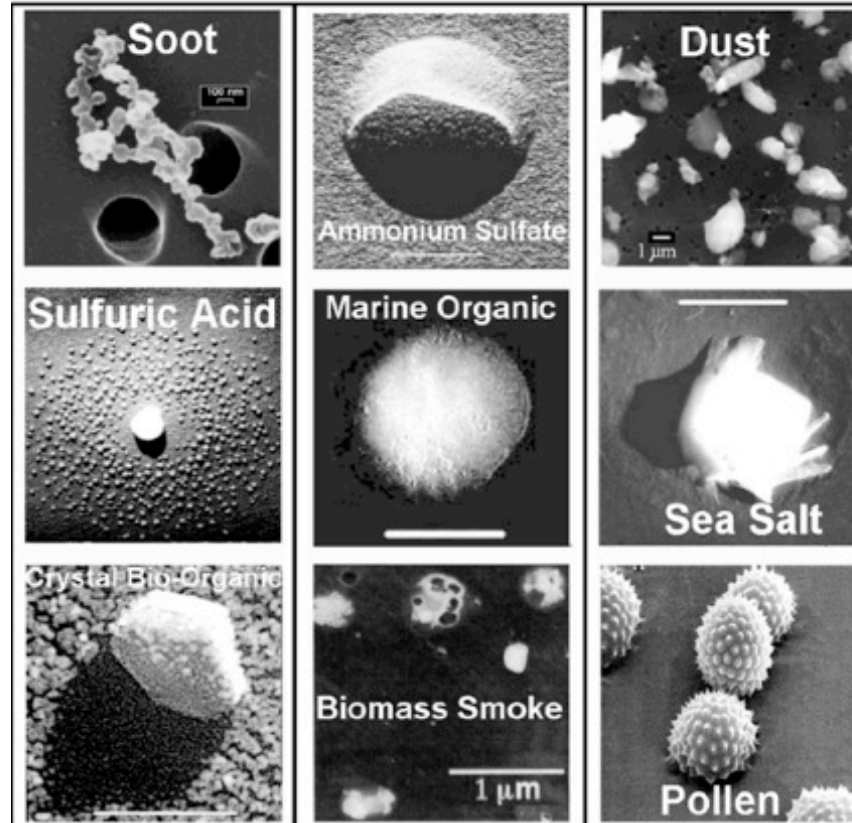
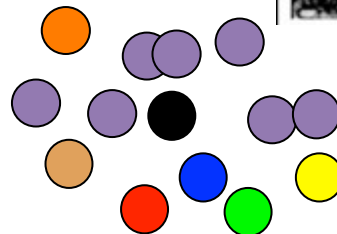
## Internal Mixture:

Each particle is same composition: BC+sulfate+sea salt



## Reality:

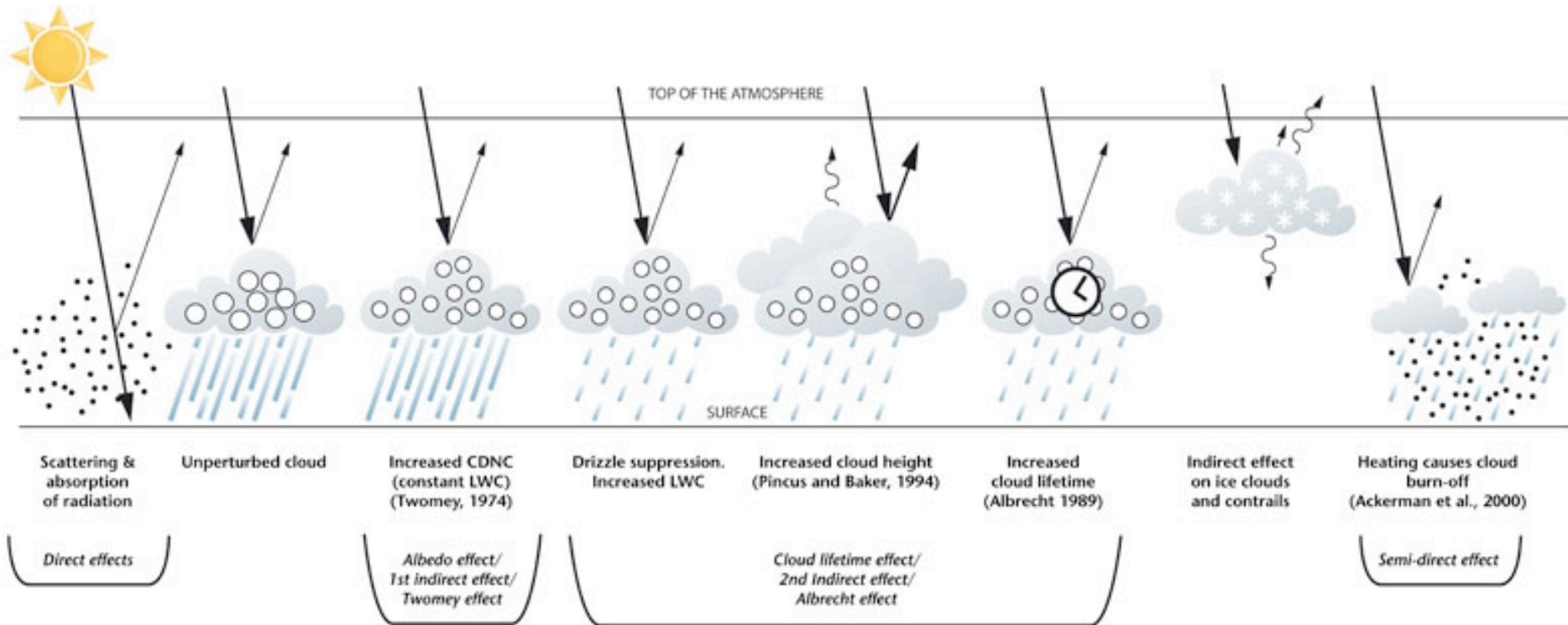
Have aerosols that are both external and internal mixtures



Composition is important in affecting radiation – extent of scattering vs absorbing

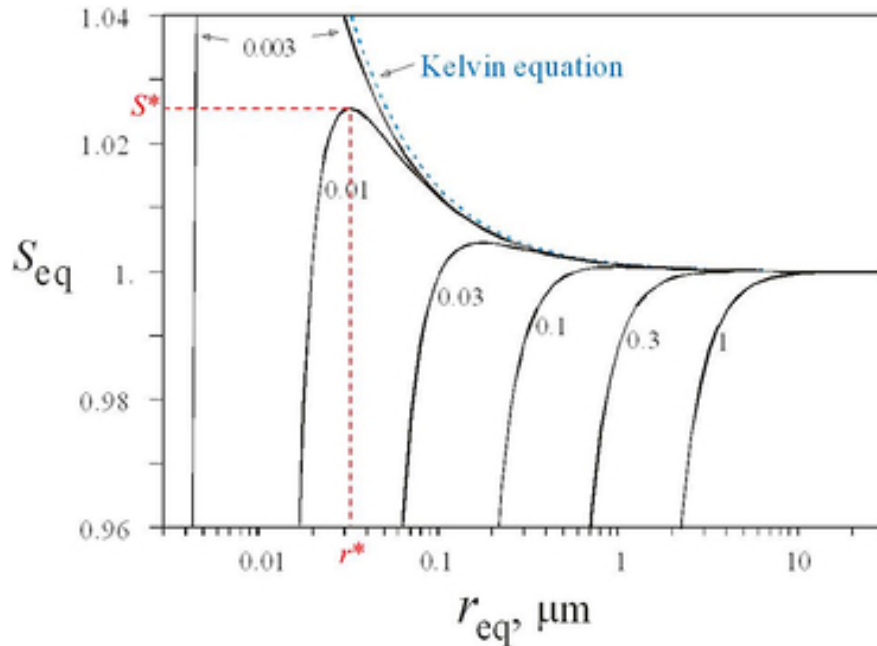
# Aerosol Effects on Clouds and Radiation

## Background





# Aerosols and Clouds



Cloud drops form on aerosols:  
Cloud Condensation Nuclei (CCN)

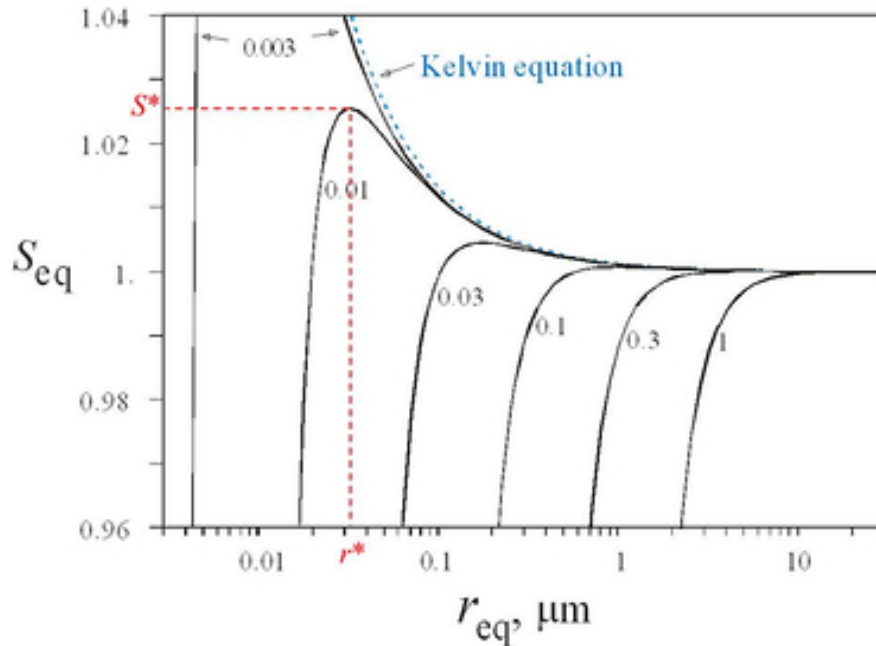
Köhler curves show relation between the equilibrium saturation ratio over the solution drop surface as a function of drop radius

Drops contain dissolved salt causing a solute effect (Raoult) reducing surface vapor pressure

Curvature effect (Kelvin) increases saturation ratio

Peak saturation ratio at a unique critical radius

# Aerosols and Clouds



## Important Factors for Drop Formation:

**Supersaturation** ( $RH > 100\%$ )

**Particle Size**

**Particle Chemical Composition**

-- alters Raoult (via water activity) and Kelvin effects (via surface tension)

**Temperature**

-- aerosol temperature (e.g. warmed by latent heating, or solar radiation absorption) affects saturation vapor pressure, surface tension, solution activity

How is this represented in cloud-resolving to global models?

# Aerosols and Clouds

$$N_{CCN} = C S^k$$

- Simple and effective formulation
- But  $N_{CCN}$  is always increasing – unrealistic
- No dependence on chemical composition

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# Aerosols and Clouds

- Köhler theory to find supersaturation
- Hygroscopicity parameter,  $\kappa$ , used to represent the aerosol population's affinity for water and how readily the aerosols will activate to cloud drops

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**Particle Size**

**Particle Chemical Composition**

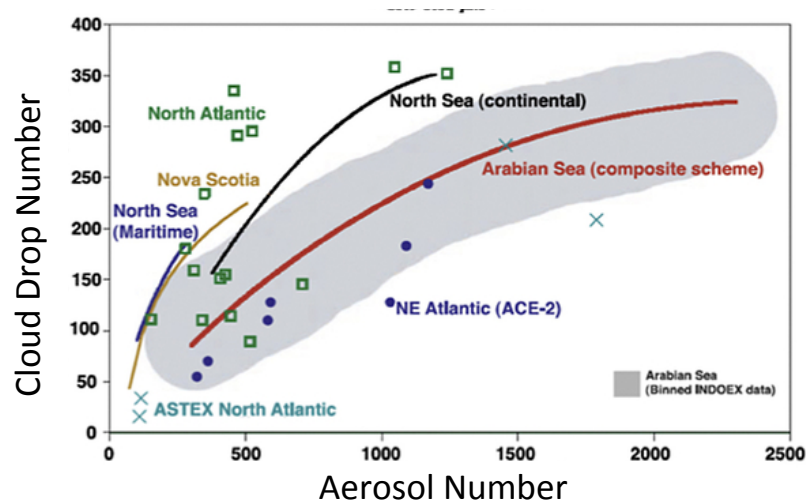
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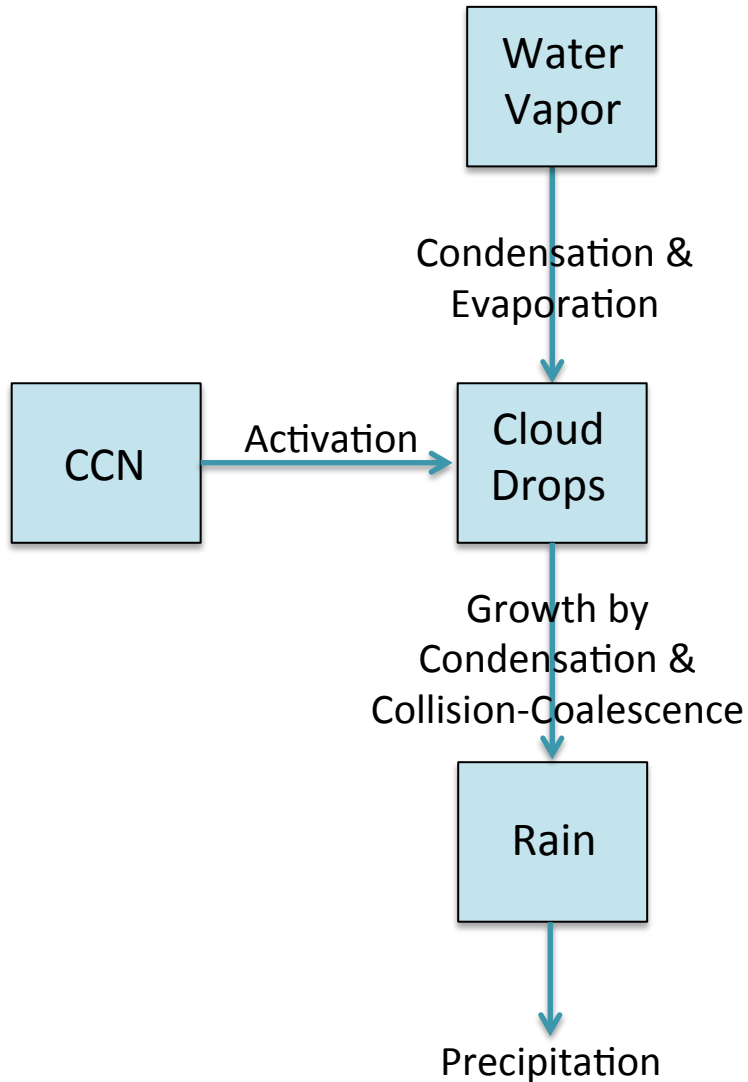
# Modeling Cloud Drop Formation

- Many cloud resolving models (CRMs) predict supersaturation  $S$  using Köhler theory with the hygroscopicity parameter  $\kappa$
- Many global models and regional cloud models use a saturation adjustment assumption and do not predict  $S$  explicitly
- Because peak supersaturation occurs within the first 50 m of cloud base, many CRMs under-predict supersaturation
- One solution is embedding Lagrangian parcel model or using a look-up table produced from a Lagrangian parcel model



Observations of  $N_{CN}$  and  $N_d$  summarized by *Ramanathan et al. (2001) Science*

# Warm Rain Formation

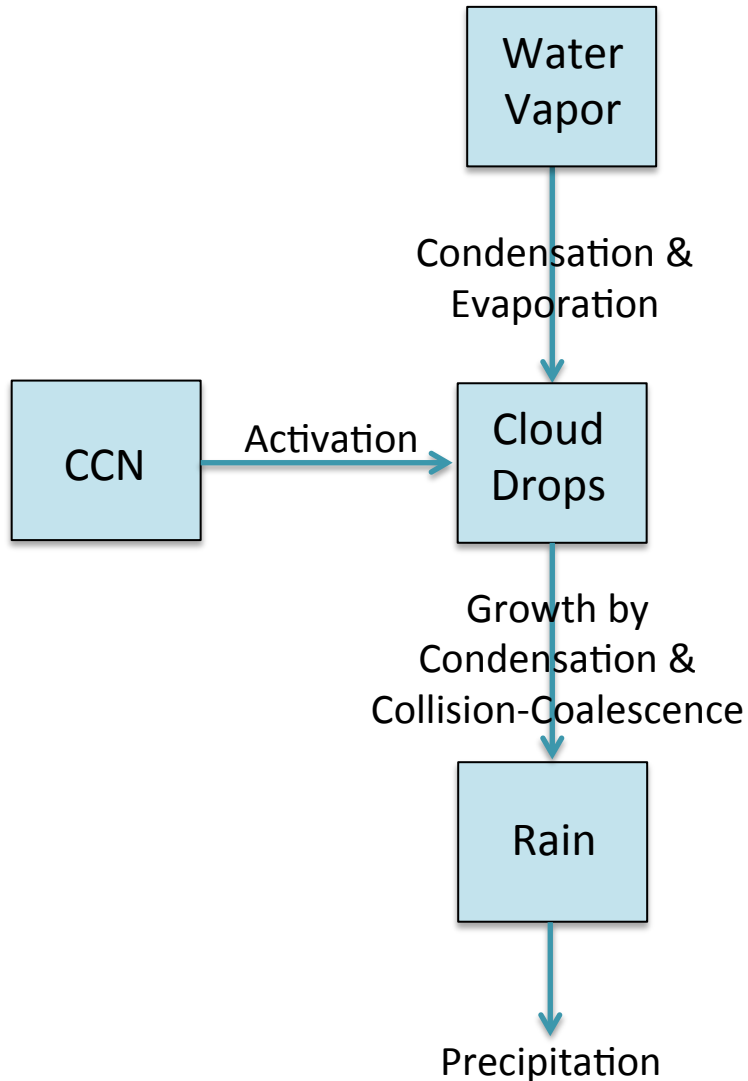


Rain formed by

- growing cloud drops,
- collision-coalescence of cloud drops,
- collection of cloud drops by rain drops



# Warm Rain Formation

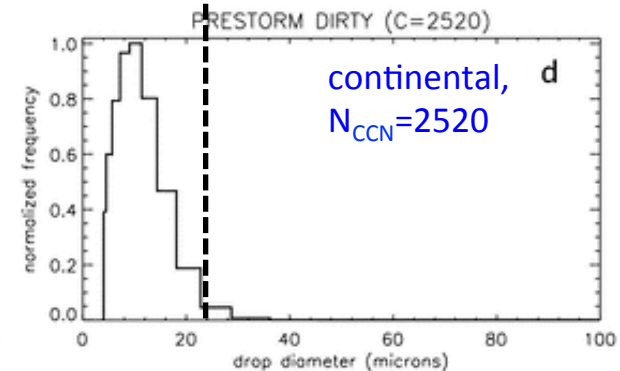
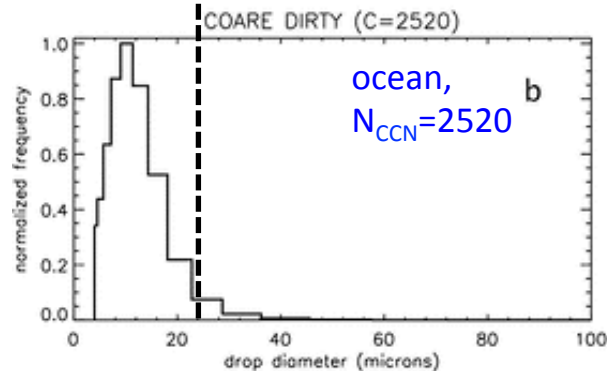
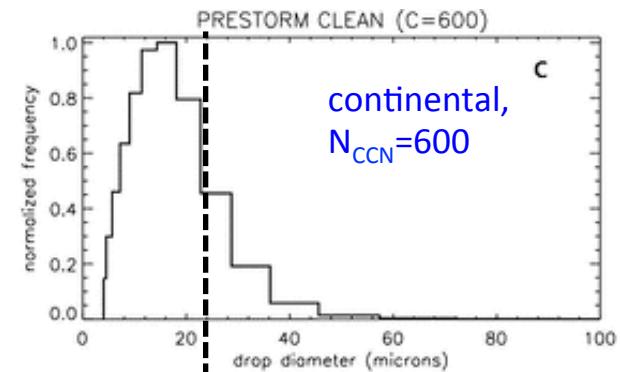
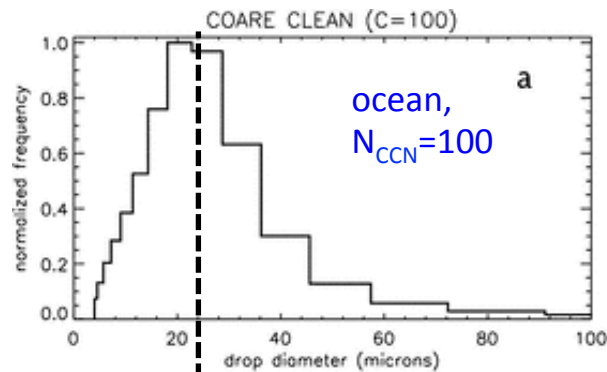


- When  $N_{\text{CCN}}$  increase,  $N_{\text{drop}}$  increase
- Higher  $N_{\text{drop}}$  leads to smaller cloud drops  $\rightarrow$  collision-coalescence becomes less efficient  $\rightarrow$  difficult to form rain (i.e., drops  $> 24 \mu\text{m}$  diameter)
- Higher  $N_{\text{drop}}$  leads to narrow cloud drop size spectrum  $\rightarrow$  less difference in fall speeds leads to suppression of rain

# Warm Rain Formation

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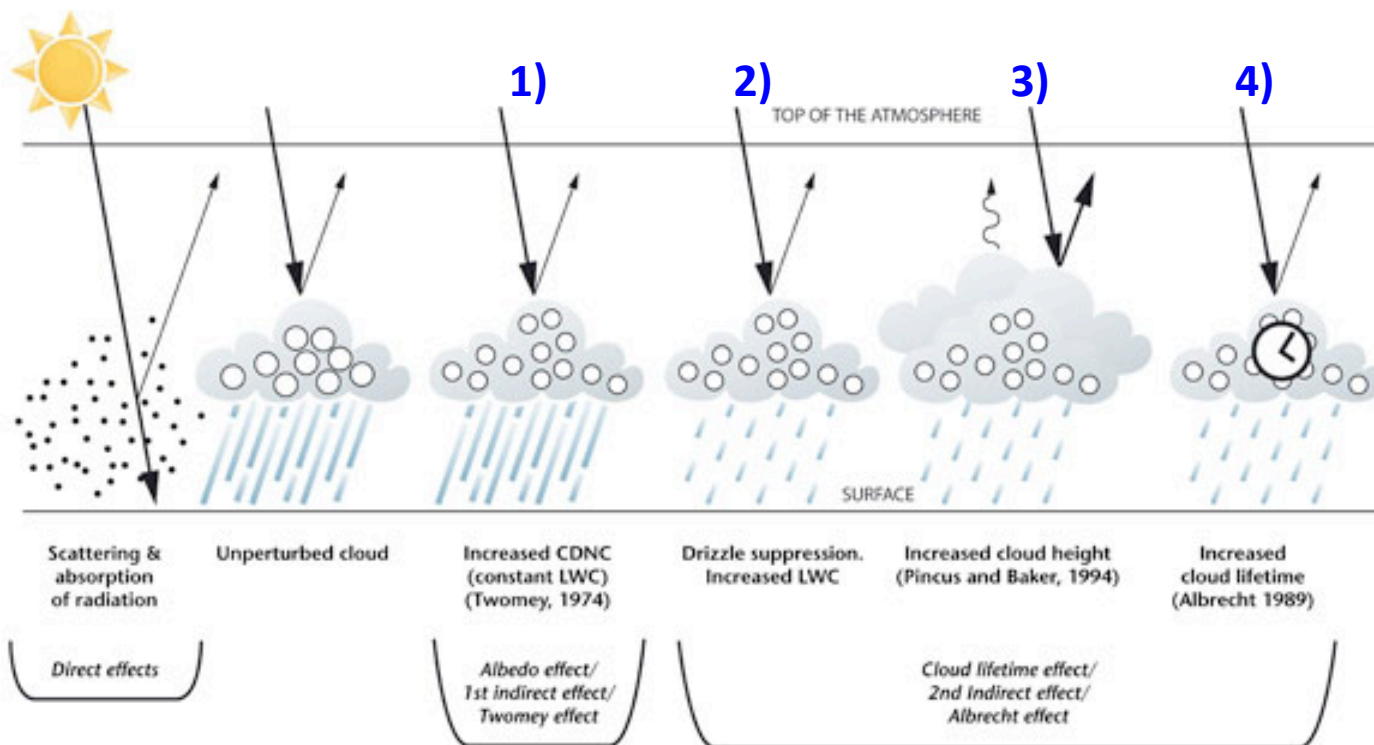
Compare drop size distributions in top panel to bottom panel



# Aerosol Effects on Clouds and Radiation

Increased number of aerosols with same liquid water content:

- 1) → increased cloud drop number and cloud albedo
- 2) → less drizzle, increased LWC
- 3) → increased cloud height
- 4) → increased cloud lifetime



# Ice Formation

Homogeneous Ice Nucleation: freezing of aerosol or cloud drops

- Occurs at  $T < -40^{\circ}\text{C}$
- Number of ice crystals controlled by number of CN entering clouds

Heterogeneous Ice Nucleation: Ice nuclei aids freezing

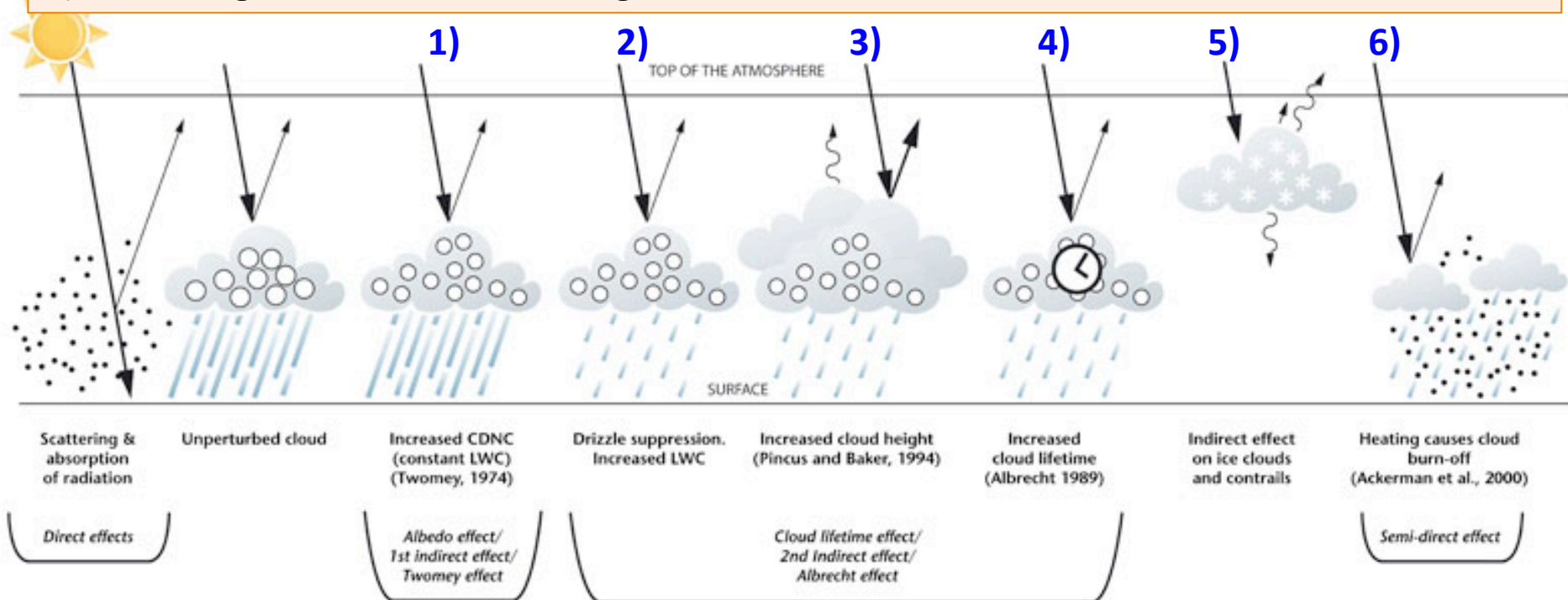
- Occurs at  $T > -40^{\circ}\text{C}$
- Water vapor directly depositing onto ice nuclei
- Cloud drops collide with aerosol at below freezing temperatures
- Cloud drops collected by ice, snow, graupel, or hail (riming)
  
- $N_{\text{CN}}$  affects ice nucleation because it affects  $N_d$
- CN affect growth of ice particles via Wegener-Bergeron-Findeisen process (water evaporating from cloud drops in order to deposit on ice)  $\rightarrow$  higher  $N_d$  causes faster WBF process
- Riming is often suppressed with more CN (similar to collision-coalescence)

Changes in the cloud physics affect the latent heating and therefore dynamics of the storm

# Aerosol Effects on Clouds and Radiation

Increased number of aerosols with same liquid water content:

- 1) → increased cloud drop number and cloud albedo
- 2) → less drizzle, increased LWC
- 3) → increased cloud height
- 4) → increased cloud lifetime
- 5) Increased number of aerosols affect cloud ice number and radiation
- 6) Absorbing aerosols cause heating → burn-off of clouds

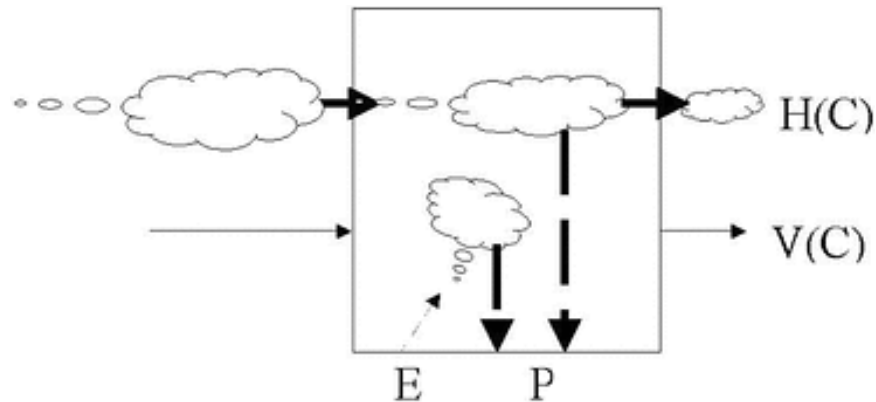




# Observing Aerosol Effects on Clouds and Radiation

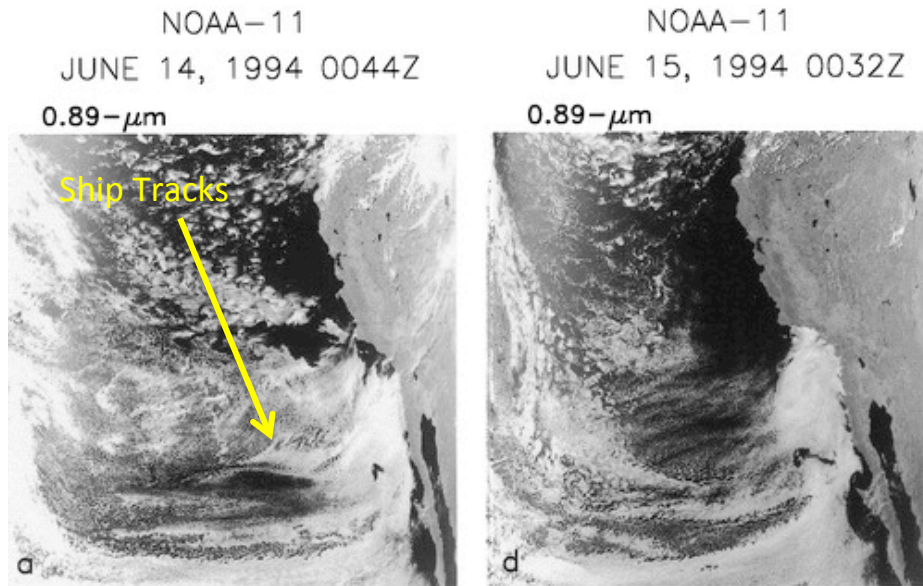
## Challenge of decoupling influences of atmospheric dynamics and thermodynamics

Major water processes occurring in an atmospheric column



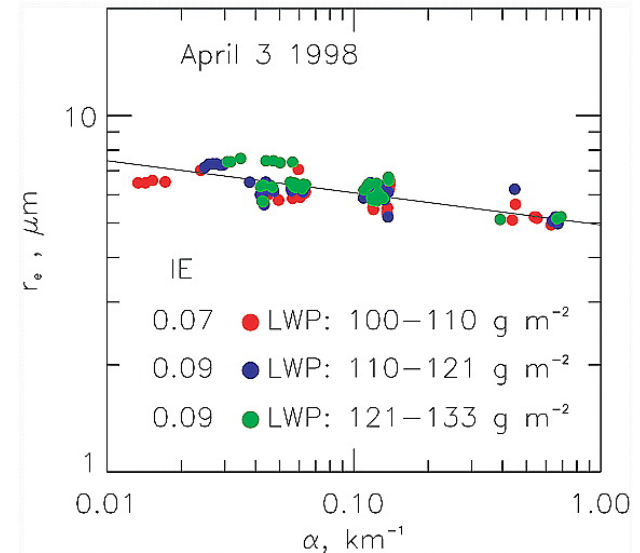
- Observations have snapshots of aerosols, clouds, precipitation
- Satellite retrievals cannot see aerosols below clouds
- What is seen in the snapshot is a result of processes over a larger spatial-temporal domain (advection of cloud, water vapor, etc)
- Extreme cases enhance the aerosol-cloud signal: ship tracks, smoke, heavy air pollution, dust storms
- Time periods when aerosol loading goes up and down along with cloud properties
- Large ensemble of cases suppress the dynamics signal

# Observing Aerosol Effects on Clouds and Radiation



Coakley et al. (2000) *J. Atmos. Sci.* -

Cloud drop effective radius from ground-based remote sensors and aerosol extinction coefficient from Raman lidar.



Feingold et al. (2003) *Geophys. Res. Lett.*

- Extreme cases enhance the aerosol-cloud signal: ship tracks, smoke, heavy air pollution, dust storms
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# Aerosol Effects on Clouds and Precipitation

- Do increased aerosol number concentrations increase or decrease precipitation in storms?
  - In convective clouds?
  - Severe convection? hurricanes, tornadoes
- Combining cloud physics and storm dynamics

# Aerosol Invigoration Effect

Changes in the cloud physics affect the latent heating and therefore dynamics of the storm

# Aerosol Invigoration Effect

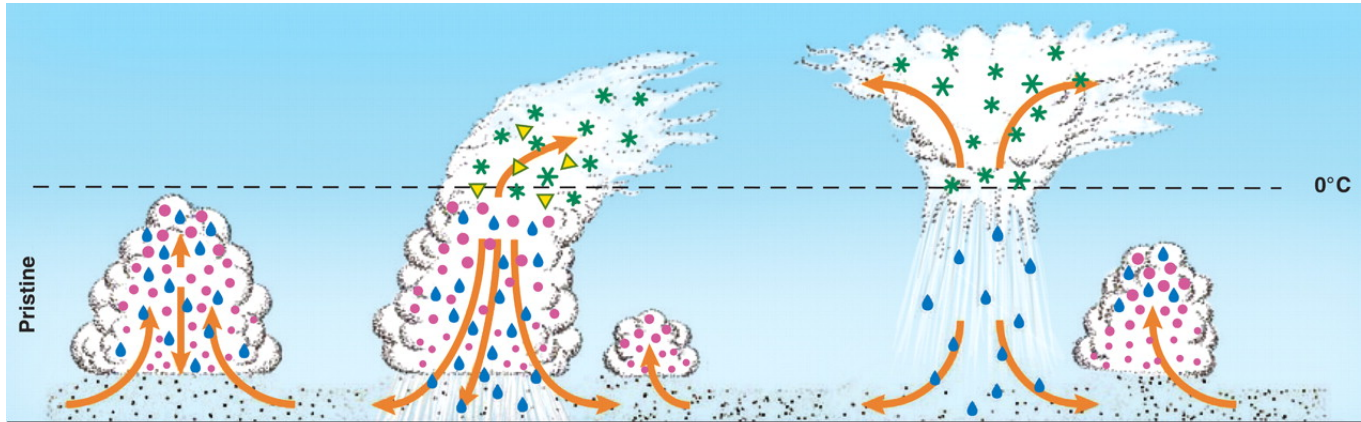
Changes in the cloud physics affect the latent heating and therefore dynamics of the storm

- More CCN → more cloud droplets, suppresses coalescence and warm rain formation
- Insufficient growth below the freezing level allows more cloud drops to be lofted above the freezing level (into the mixed phase region)
- Evidence of sustained supercooled liquid water at  $T \geq -37.5^{\circ}\text{C}$  in continental convection
- More freezing of cloud drops and associated latent heat release, enhances the growth of large hail and cold-rain processes
- Onset of precipitation is delayed; model results suggest greater intensity of rain later in life cycle of cloud
- Latent heat release increases updraft speeds
- Melting and evaporation cause more cooling; enhancing the cold pool

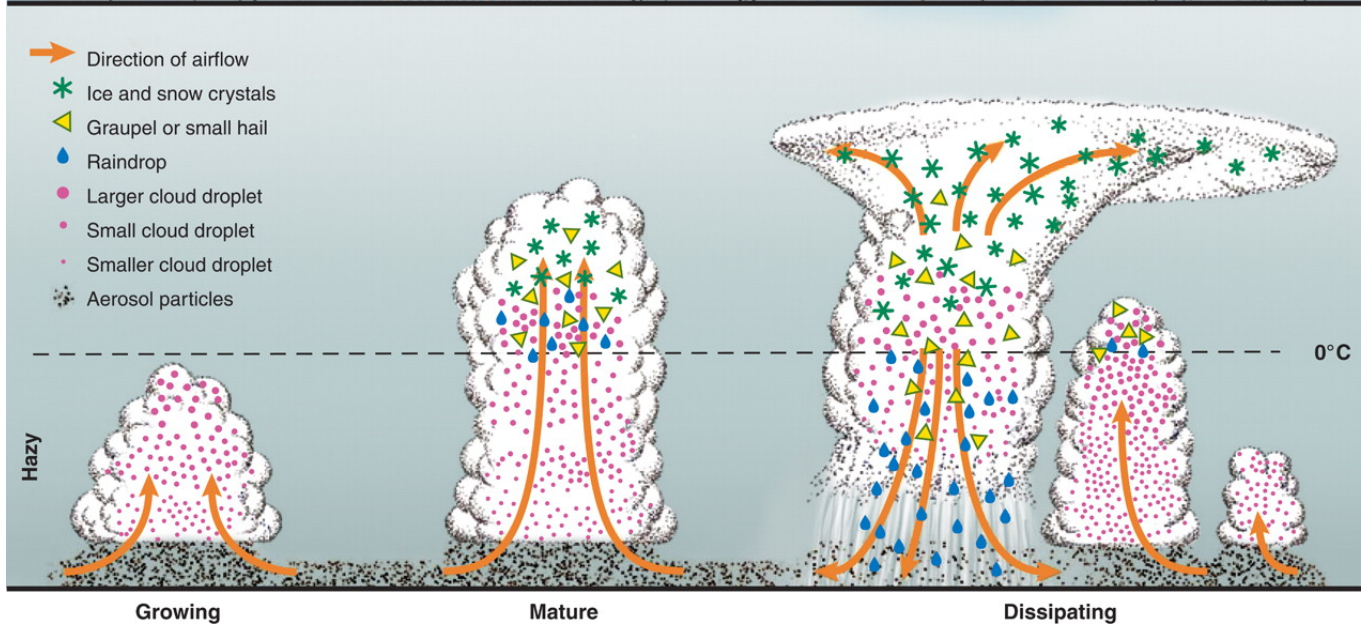


# Aerosol Invigoration Effect

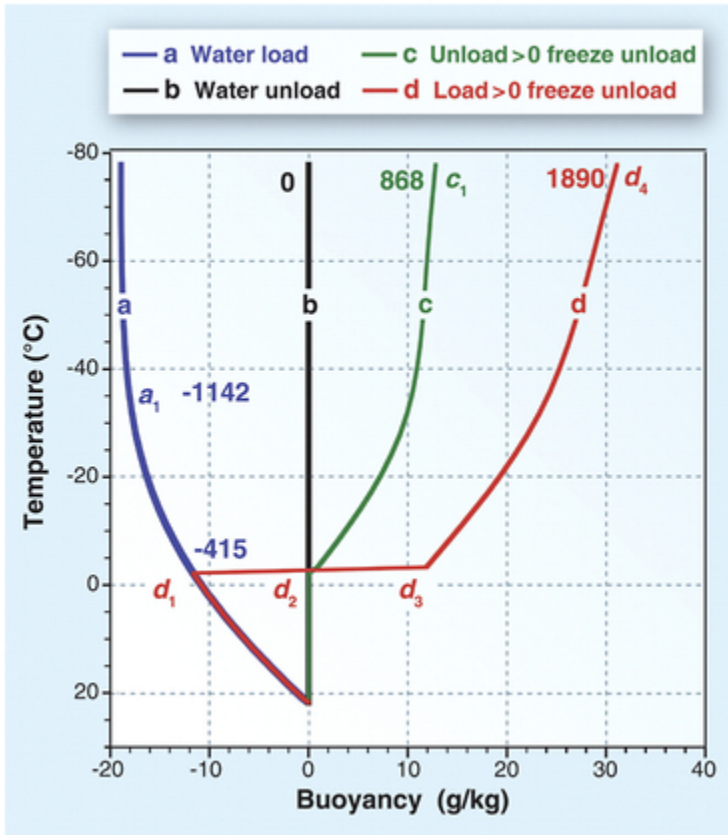
Low aerosol number concentrations



High aerosol number concentrations



# Aerosol Invigoration Effect – buoyancy



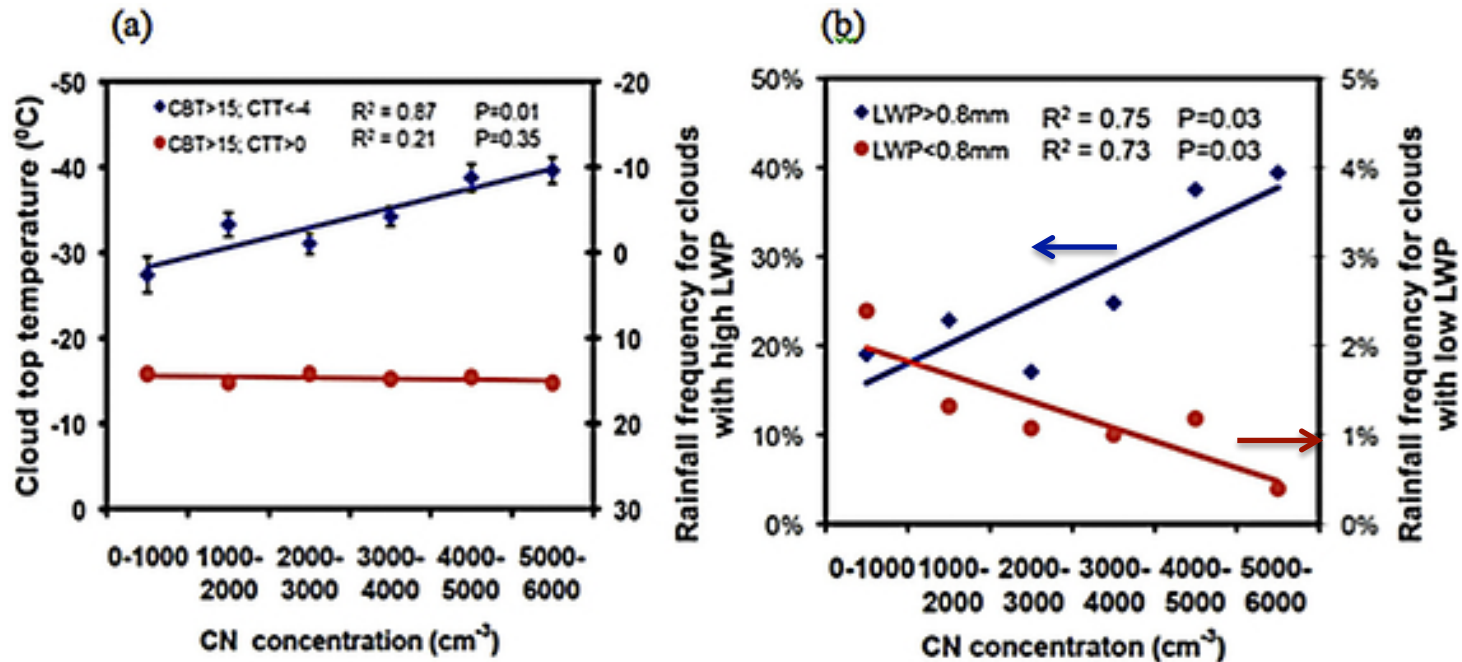
- Suppressing rainfall and keeping all condensed water without freezing
- Precipitating all condensed water without freezing
- Precipitating all condensates with freezing at  $T < -4^\circ\text{C}$
- Suppressing precipitation until  $T = -4^\circ\text{C}$  and then freezing and precipitating all condensed water

Released static energy (J/kg) with respect to scenario b is given for a, c, d

Latent heat release increases updraft speeds

# Evidence of Aerosol Invigoration Effect

## 10 years of data from DOE ARM Southern Great Plains site



- As CN increases cloud top temperature decreases (i.e. cloud top height increases) when cloud top temperature is < -4°C
- No effect for warmer clouds
- As CN increases frequency of rainfall increases when LWP > 0.8 mm
- As CN increases frequency of rainfall decreases when LWP < 0.8 mm

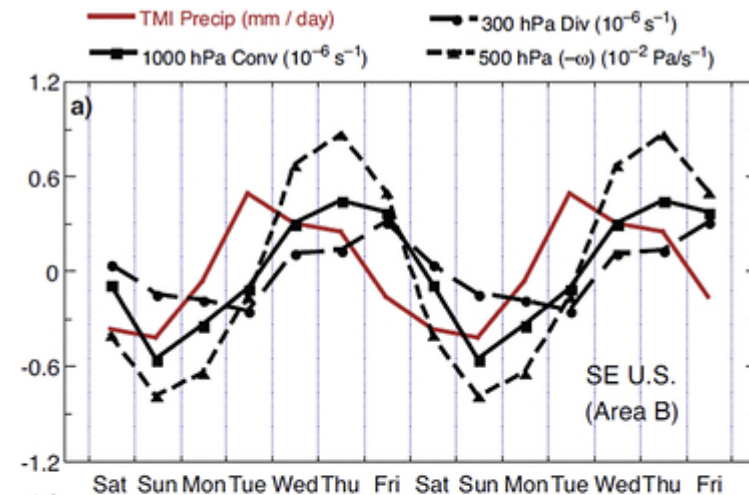
10 years of data from DOE ARM Southern Great Plains site

# Role of Warm Cloud Depth

- For higher cloud bases, the warm cloud depth is shorter and there will not be very much liquid water to release latent heat
  - For lower cloud bases, the warm cloud depth is deeper and the liquid water has time to grow via collision-coalescence causing drops to rain out before lofting above the freezing level
- Aerosol invigoration occurs in regions with the right environmental conditions

# Do Aerosols from Urban Centers Affect Precipitation?

- High concentrations of hygroscopic aerosols
- Warmer temperatures in urban center compared to outside urban center
- Moist, hot air providing high CAPE
- Satellite data analysis points to a weekly cycle in precipitation in summer over the southeastern U.S.
- Controversial result – other studies show lack of weekly cycle or need for better statistical analysis (Schultz et al., 2008; Murphy et al., 2008; Kim et al., 2010; Tuttle and Carbone, 2011)





# Absorbing Aerosols and Clouds

- Thermodynamics of the atmosphere is altered through heating by the absorbing aerosols (black carbon, smoke)
- Depending on the height of the aerosols, the atmospheric stability is enhanced, reducing moisture content from evaporation and inhibiting formation of clouds and precipitation (Ackerman et al., 2000)
- Smoke aerosols affect clouds via CCN and via heating – two competing processes causing increases in cloud top height for aerosol optical depth  $< 0.4$  and decreases in cloud top height at higher AOD

# Modeling Aerosol Effects on Clouds and Precipitation

